

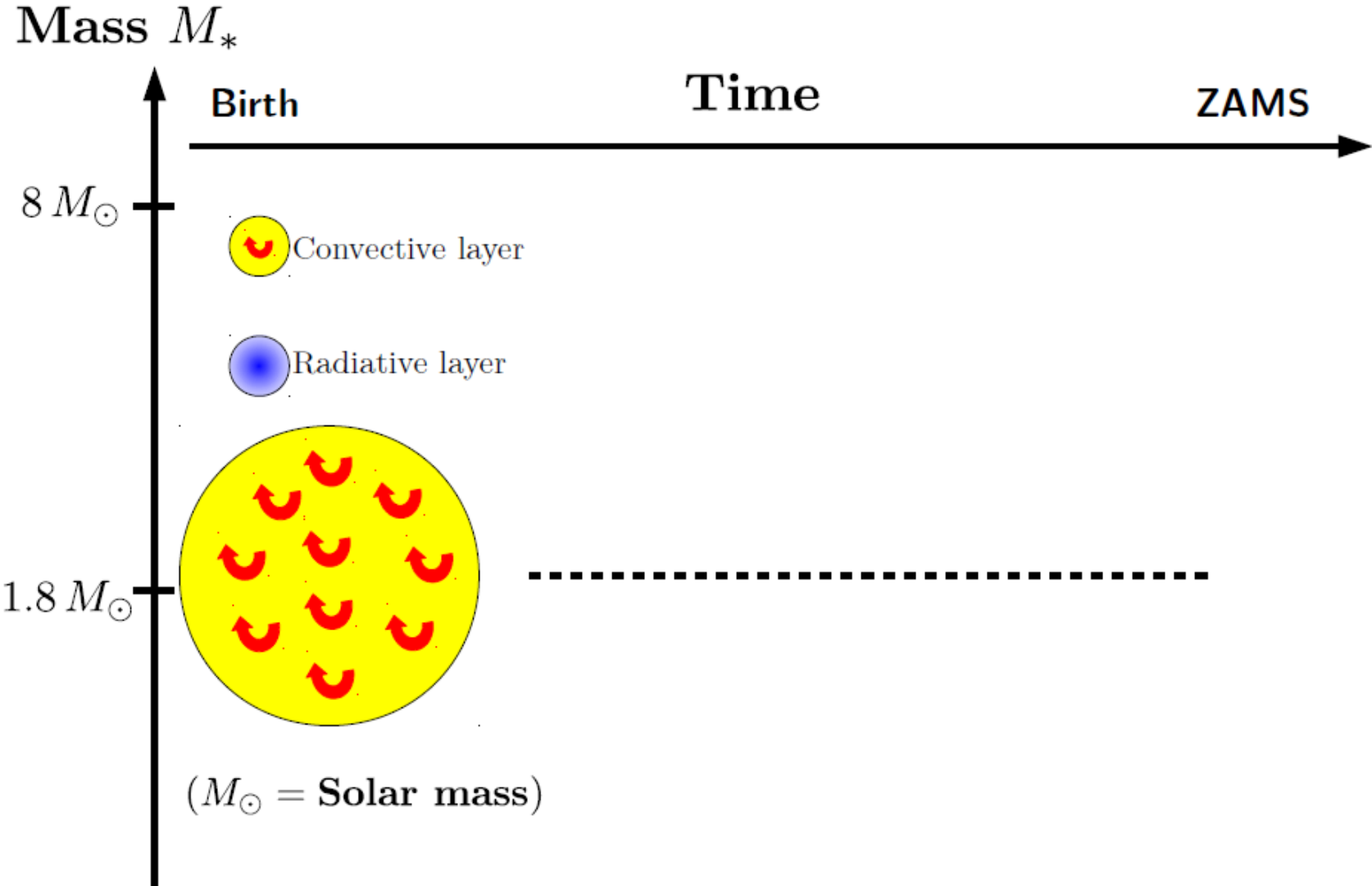
Magnetic fields driven by tidal mixing in radiative stars

J. Vidal¹, David Cébron¹, R. Hollerbach², N. Schaeffer¹

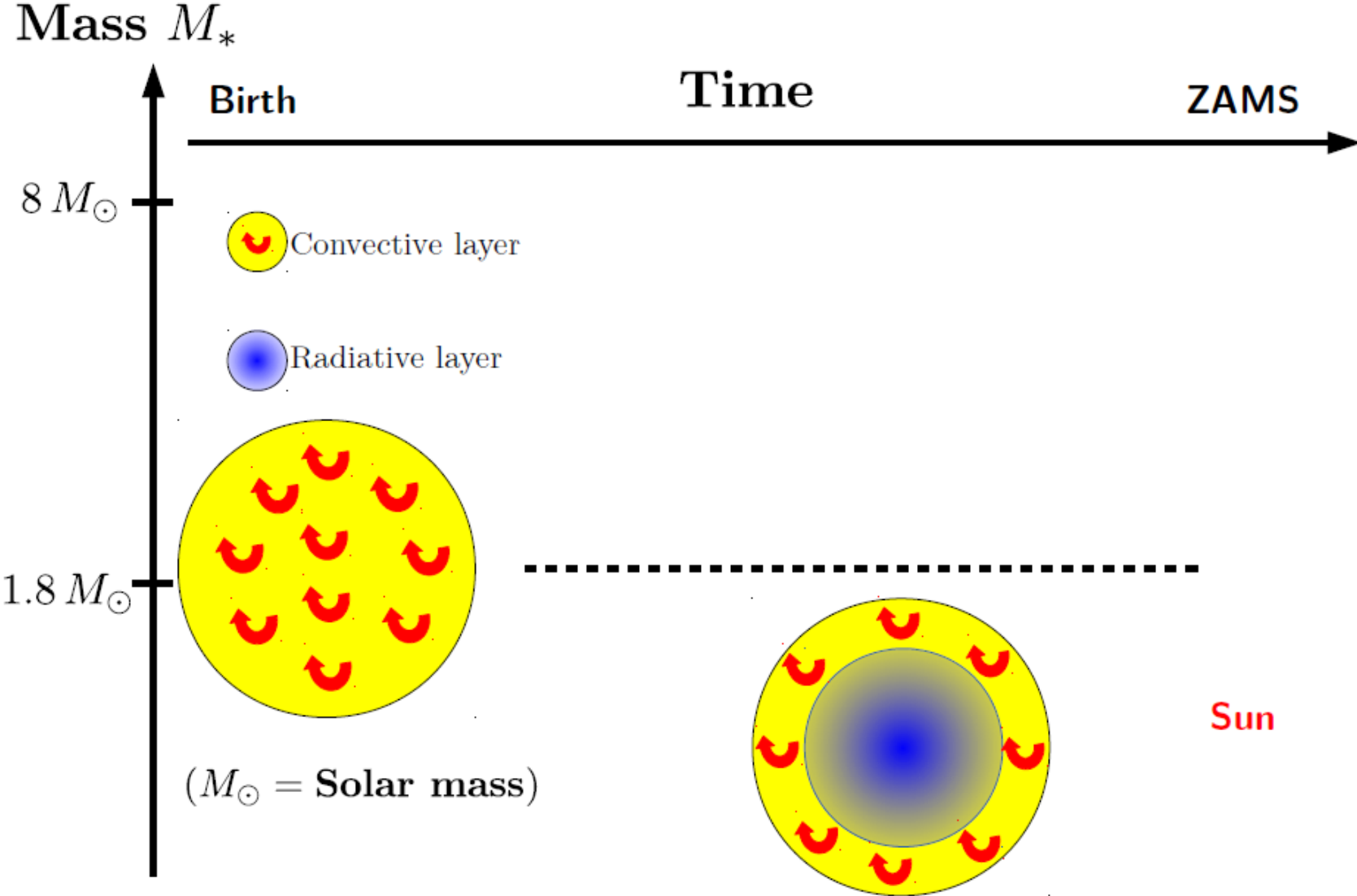
¹ *ISTerre (CNRS, Univ. Grenoble Alpes)*
² *Univ. of Leeds*

Montpellier, 28 Mars 2018

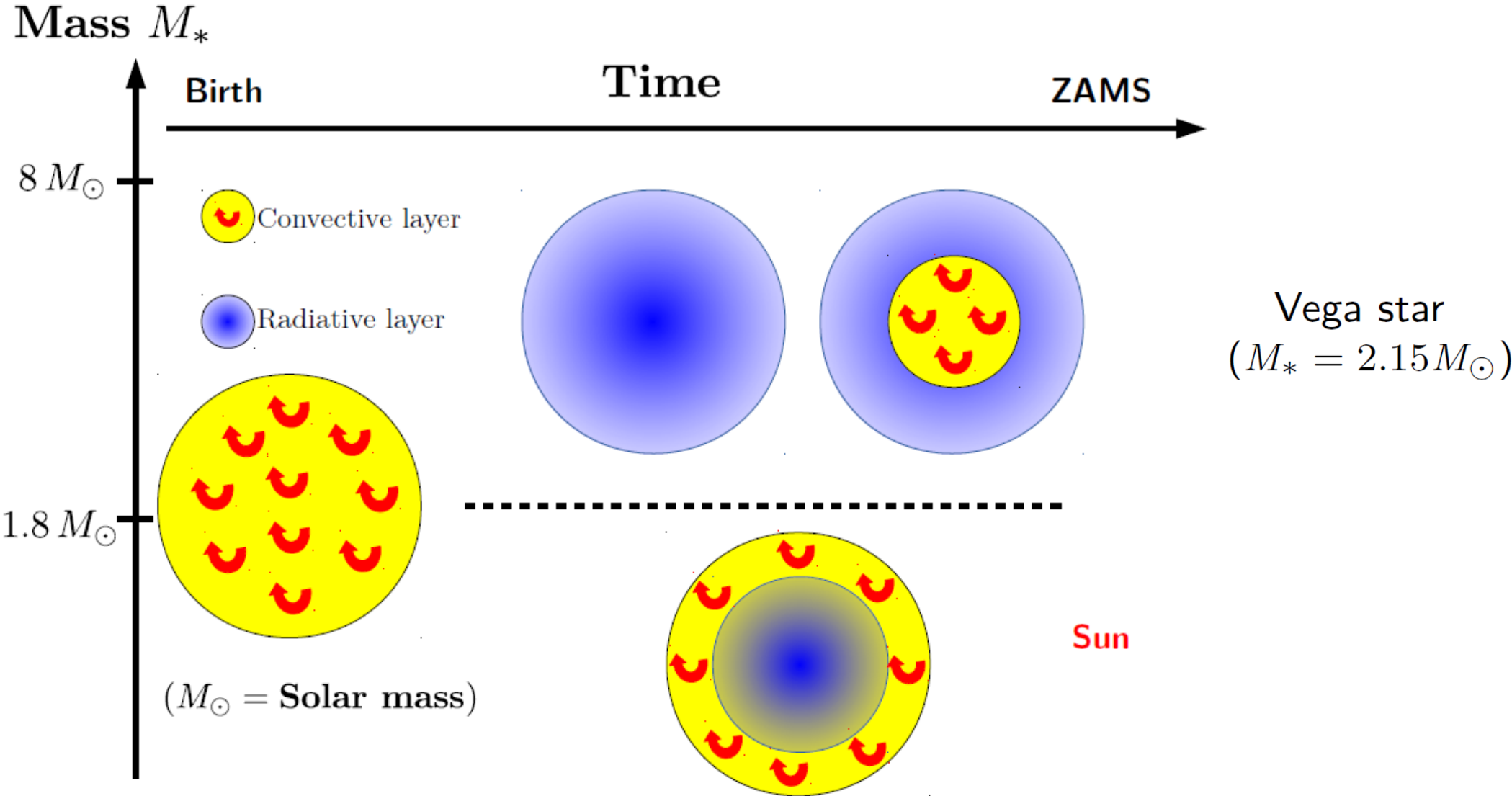
Stellar context



Stellar context



Stellar context



Magnetism of intermediate mass, radiative stars

10% with surface magnetic field B_0

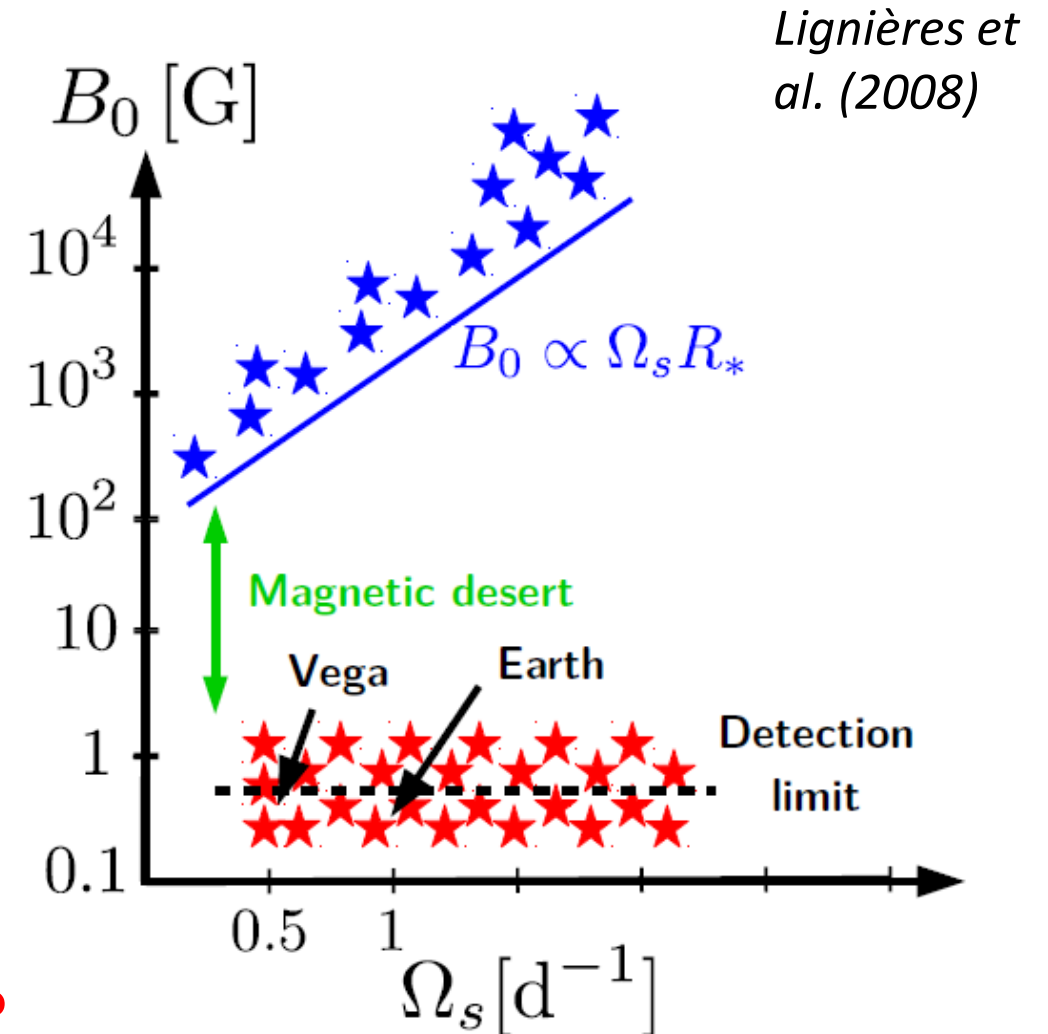
Strong fields B_0 (e.g. Ap stars)

- Typical **strength**: $10^2 - 10^4$ G
- Often **static** fields
- Surface fields B_0 with a **fossil** origin

Ultra-weak fields B_0 (e.g. Vega)

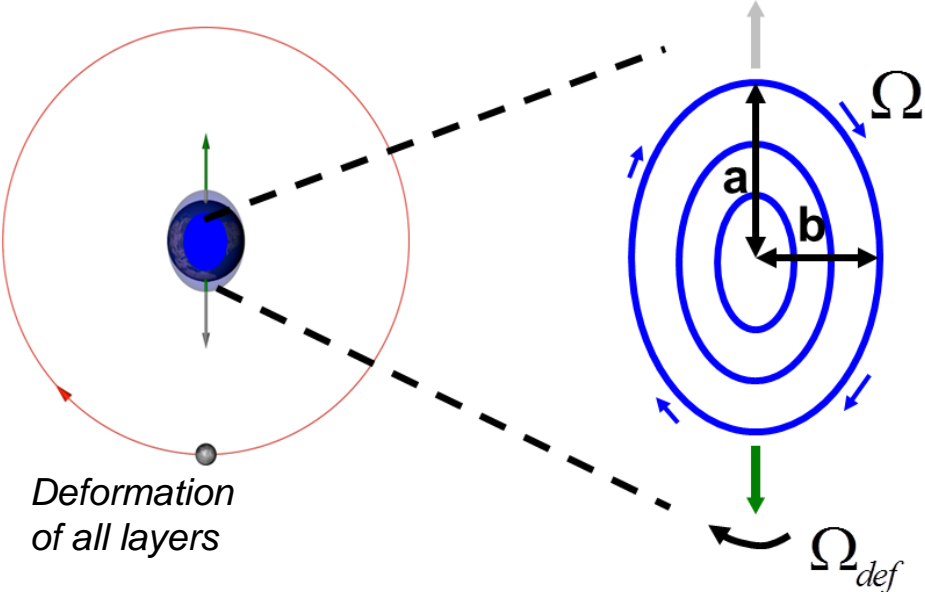
- Typical **strength**: 0.1 - 1 G
- Often **dynamical** fields
- Surface fields B_0 generated by **dynamamos**?

Tidal dynamos for Vega-like magnetism?



Tidal instability

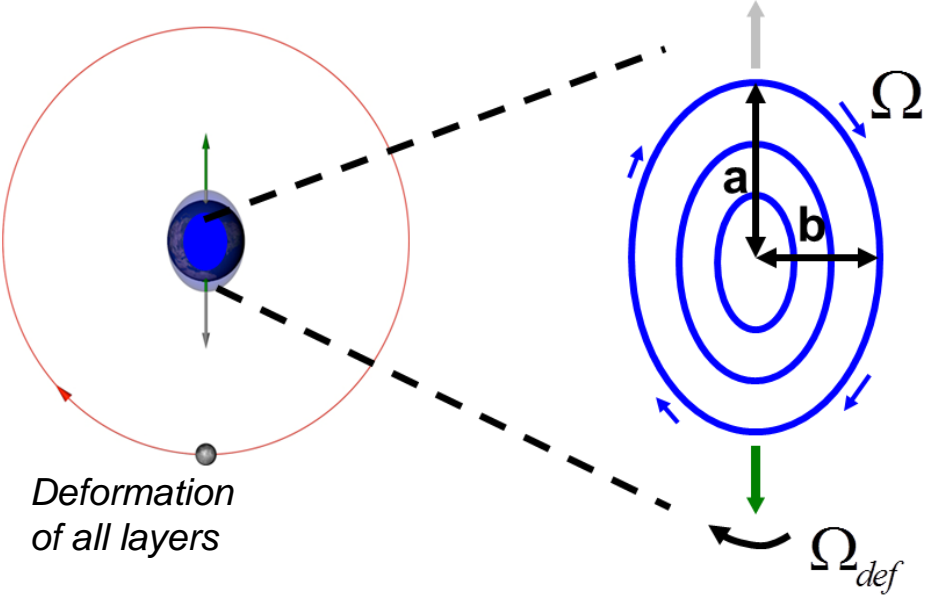
Tides → Elliptical streamlines



Deformation of all layers

Tidal instability

Tides → Elliptical streamlines



Elliptical or tidal instability!

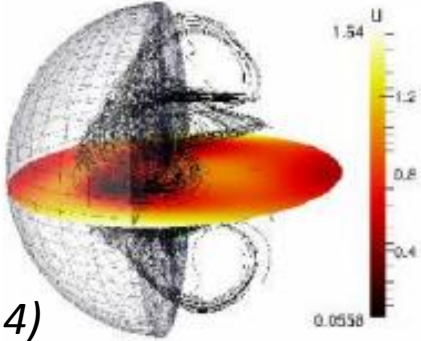
Turbulent flows



?

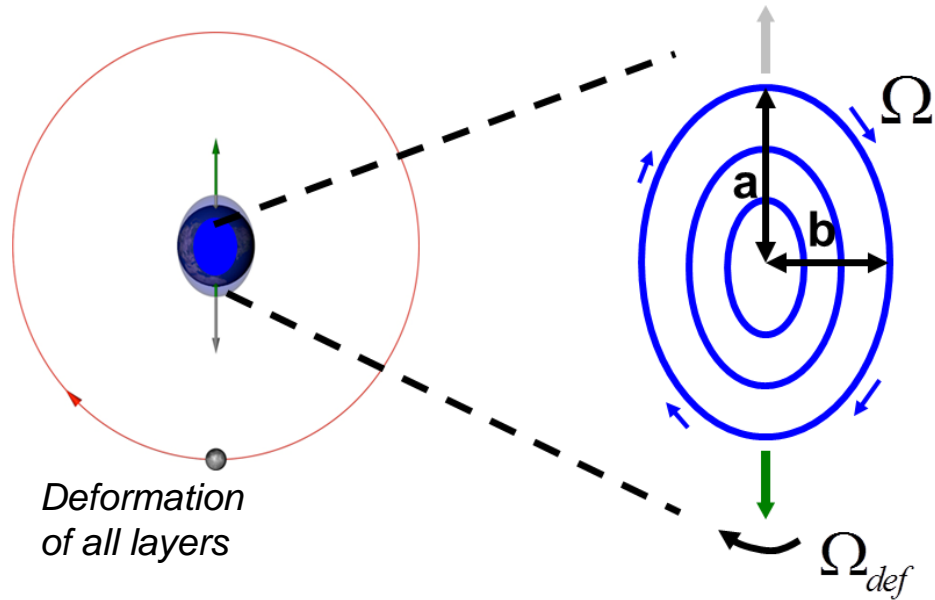
Magnetic field

Cébron & Hollerbach (2014)



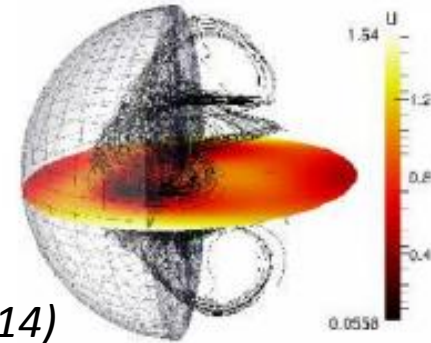
Tidal instability

Tides → Elliptical streamlines



→ Turbulent flows
Elliptical or tidal instability!

Cébron & Hollerbach (2014)



↓ ?
Magnetic field

- **OK for neutral or convective fluids** (Cébron et al. 2010)
- But **tidally driven dynamos in stably stratified fluids?**
→ radiative stars, lunar core, Encelade, Earth's liquid core stratified layer, etc.

Stable stratification

- **(Dimensional) Brunt-Väisälä frequency N_0**

$$N_0^2 = -\alpha \nabla T \cdot g$$

(α : Coeff. of thermal expansion, $g = -\nabla\Phi$: gravity)

- ▶ $N_0^2 > 0 \implies$ **stably stratified**,
- ▶ **Internal** gravity waves.

Stable stratification

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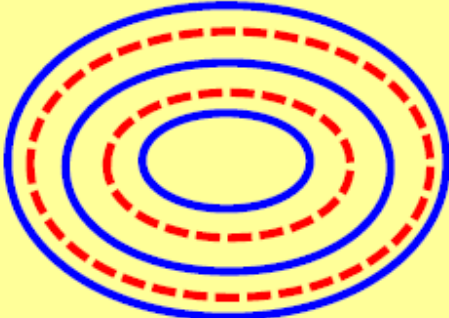
- ▶ $N_0^2 > 0 \implies$ **stably stratified**,
- ▶ **Internal** gravity waves.

- **Growth rate elsewhere? Dynamo?**

- **Hypothesis:**
 - **Boussinesq** fluid
 - **Barotropic & linear** gravity

Barotropic

T



$\nabla T \times \nabla \Phi = \mathbf{0}$

Φ ($g = -\nabla\Phi$)

- ▶ No **baroclinic** instability,
- ▶ Growth rate **reduced** at the poles & equator.

Mathematical formulation

- ▶ Basic state: U_0, T_0, g ,
- ▶ Scales: $R_*, \Omega_s^{-1}, \Omega_s^2 R_*/(\alpha g_0)$ and $R_* \Omega_s \sqrt{\mu_0 \rho_*}$,
- ▶ **Dimensionless** numbers
- ▶ Ellipsoidal container of equatorial ellipticity β_0 .

$$Ek = \frac{\nu}{\Omega_s R_*^2} \leq 10^{-16}, \quad 10^{-6} \leq Pr = \frac{\nu}{\kappa} \leq 10^{-4}, \quad 10^{-8} \leq Pm = \frac{\nu}{\eta_m} \leq 10^{-4},$$

Boussinesq equations of the perturbations (u, θ, B) in the **inertial** frame

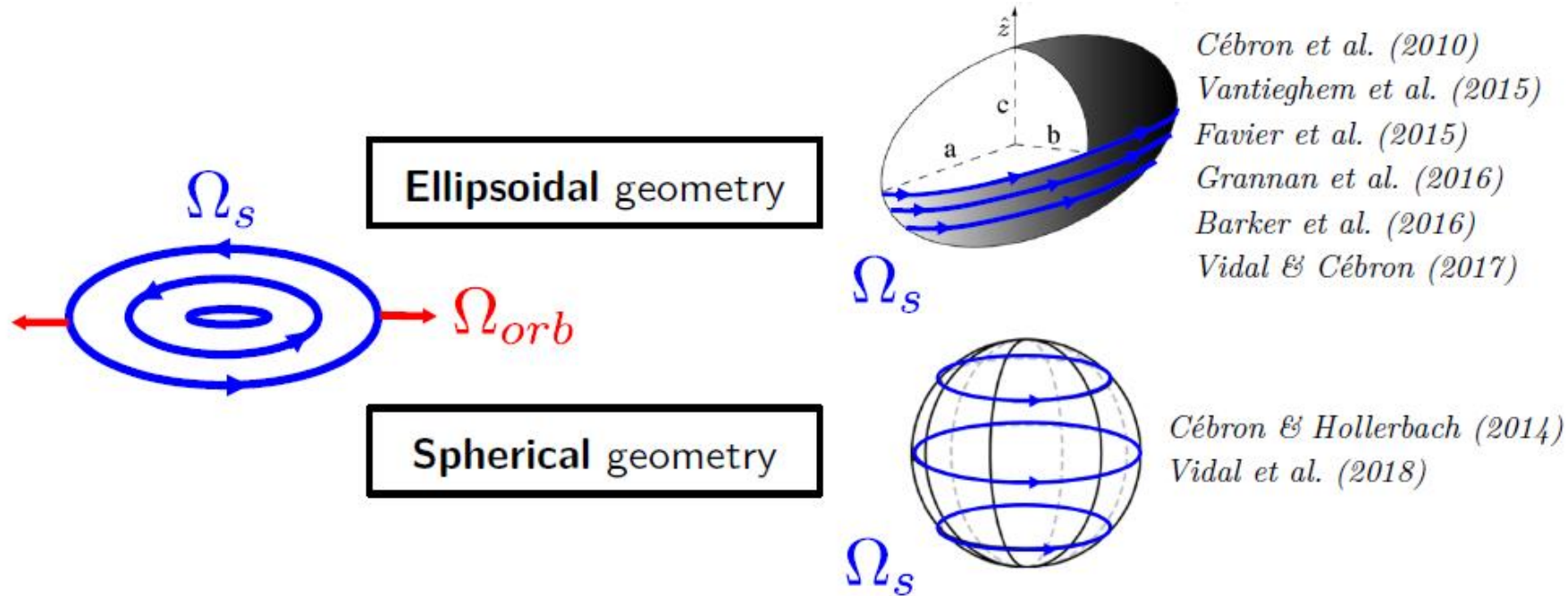
$$\frac{\partial u}{\partial t} = -(\mathbf{u} \cdot \nabla) \mathbf{U}_0 - (\mathbf{U}_0 \cdot \nabla) \mathbf{u} - (\mathbf{u} \cdot \nabla) \mathbf{u} - \nabla p + Ek \nabla^2 \mathbf{u} - \theta \mathbf{g} + (\nabla \times \mathbf{B}) \times \mathbf{B},$$

$$\frac{\partial \theta}{\partial t} = -(\mathbf{U}_0 \cdot \nabla) \theta - (\mathbf{u} \cdot \nabla) T_0 - (\mathbf{u} \cdot \nabla) \theta + \frac{Ek}{Pr} \nabla^2 \theta,$$

$$\frac{\partial \mathbf{B}}{\partial t} = \nabla \times (\mathbf{U}_0 \times \mathbf{B}) + \nabla \times (\mathbf{u} \times \mathbf{B}) + \frac{Ek}{Pm} \nabla^2 \mathbf{B},$$

$$\nabla \cdot \mathbf{u} = \nabla \cdot \mathbf{B} = 0.$$

Numerical method



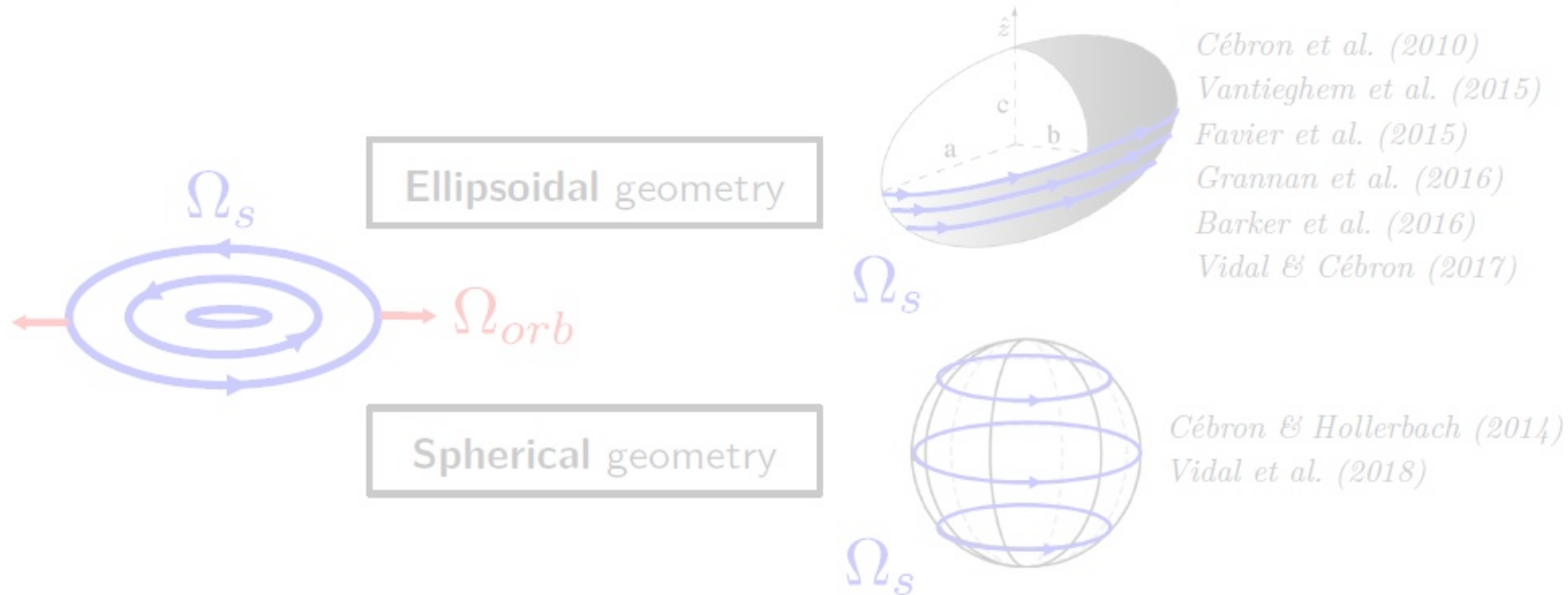
3 codes: COMSOL, YALES2, **NEK5000**

Various codes: **MAGIC**, ASH, **XSHELLS**, **PARODY**, **RAYLEIGH**

Spectral codes are more **efficient** & **reliable** for dynamo computations.

in red: open-source and freely available code

Numerical method



Cébron et al. (2010)
Vantieghem et al. (2015)
Favier et al. (2015)
Grannan et al. (2016)
Barker et al. (2016)
Vidal & Cébron (2017)

Cébron & Hollerbach (2014)
Vidal et al. (2018)

3 codes: COMSOL,
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Various codes: **MAGIC**, ASH,
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**World fastest spherical
dynamo Boussinesq code**

(benchmark of Matsui+16)

Spectral codes are more **efficient** & **reliable** for dynamo computations.

Numerical method: XSHELLS with basic states

Schaeffer (2013, 2017)

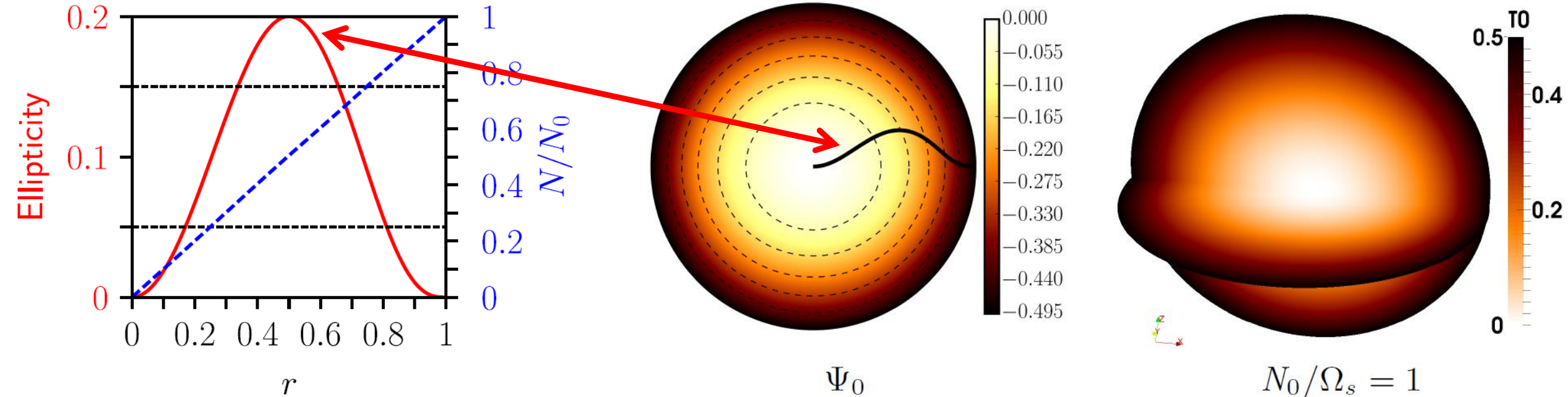
- ▶ **Open-source** code (<https://bitbucket.org/nschaeff/xshells>),
- ▶ Pseudo-spectral code with finite differences and spherical harmonics,
- ▶ Simulations **massively parallel** supercomputers (e.g. OCCIGEN): 2×10^6 h.c.

Radiative stars: idealized model (Vidal et al. MNRAS, 2018)

- ▶ Fixed $Ek = 10^{-4}$, $Pr = 1$ and $\Omega_0 = 0$,
- ▶ **Control** parameters

N_0/Ω_s and max. tidal **ellipticity** $\epsilon \ll 1$,

- ▶ **Basic** state: $U_0(\Psi_0)$, $T_0(\Psi_0)$, $g(\Psi_0)$,
- ▶ Barotropic state ($g \times \nabla T_0 = \mathbf{0}$),
- ▶ BC: **Stress-free**, **fixed** temperature and electrically **insulating**.



Hydrodynamic instability

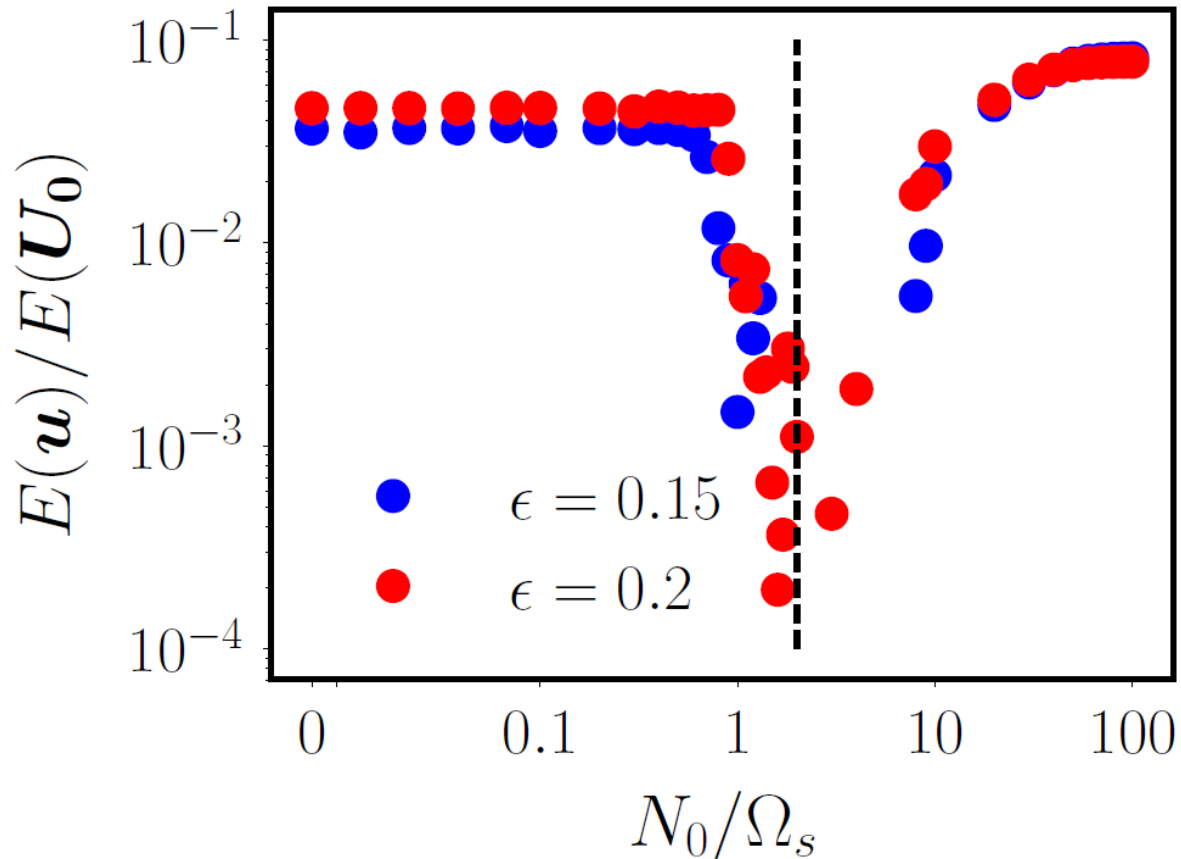
(Vidal et al. MNRAS, 2018)

- ▶ $Ek = 10^{-4}$, $Pr = 1$, $\Omega_0 = 0$,
- ▶ **Onset** at $\epsilon_c = 0.054$ ($N_0/\Omega_s = 0$)

Hydrodynamic instability

(Vidal et al. MNRAS, 2018)

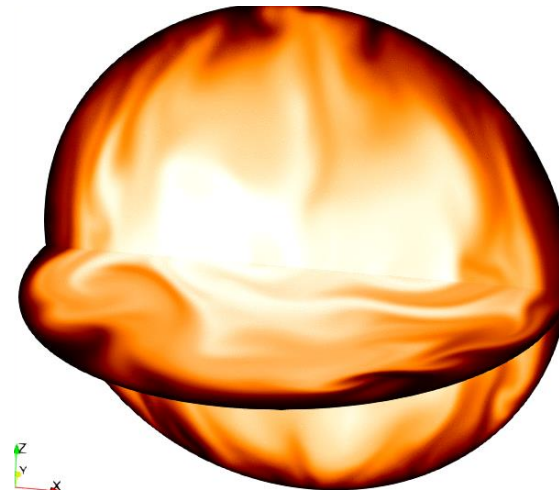
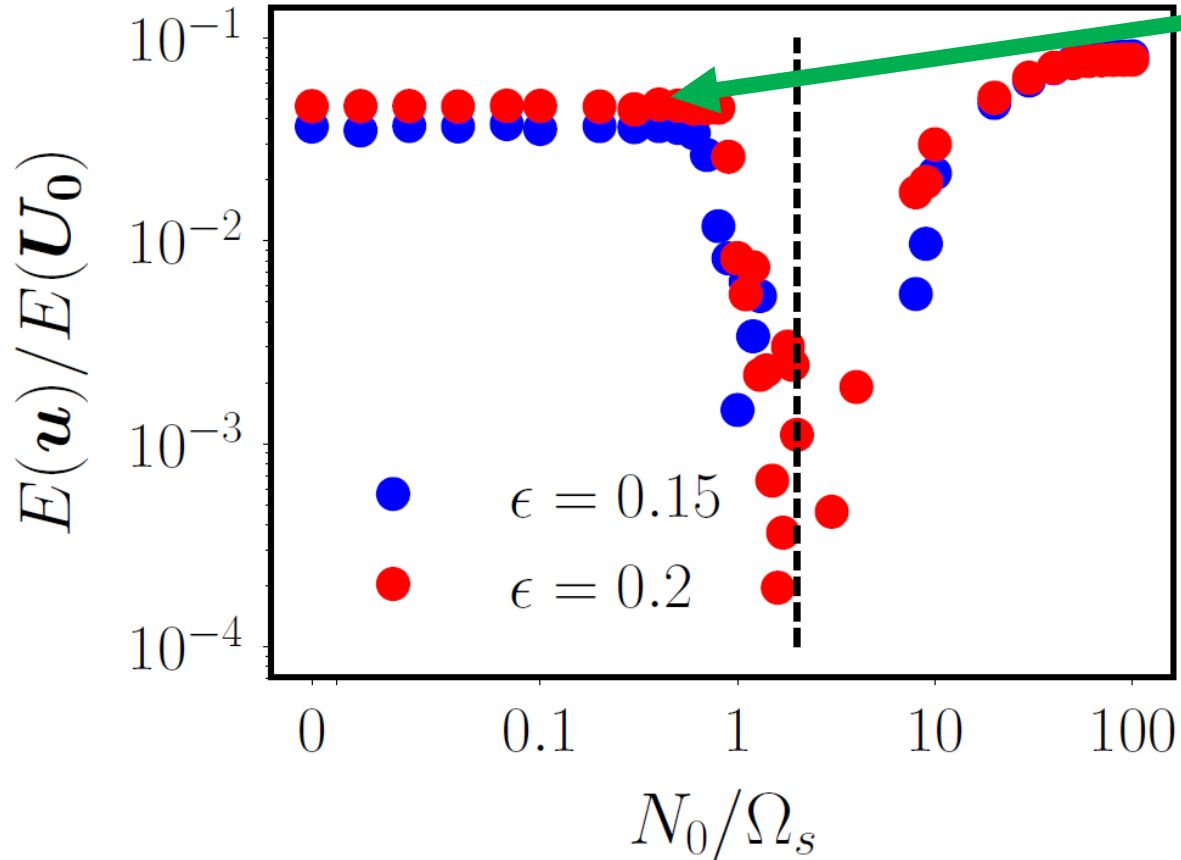
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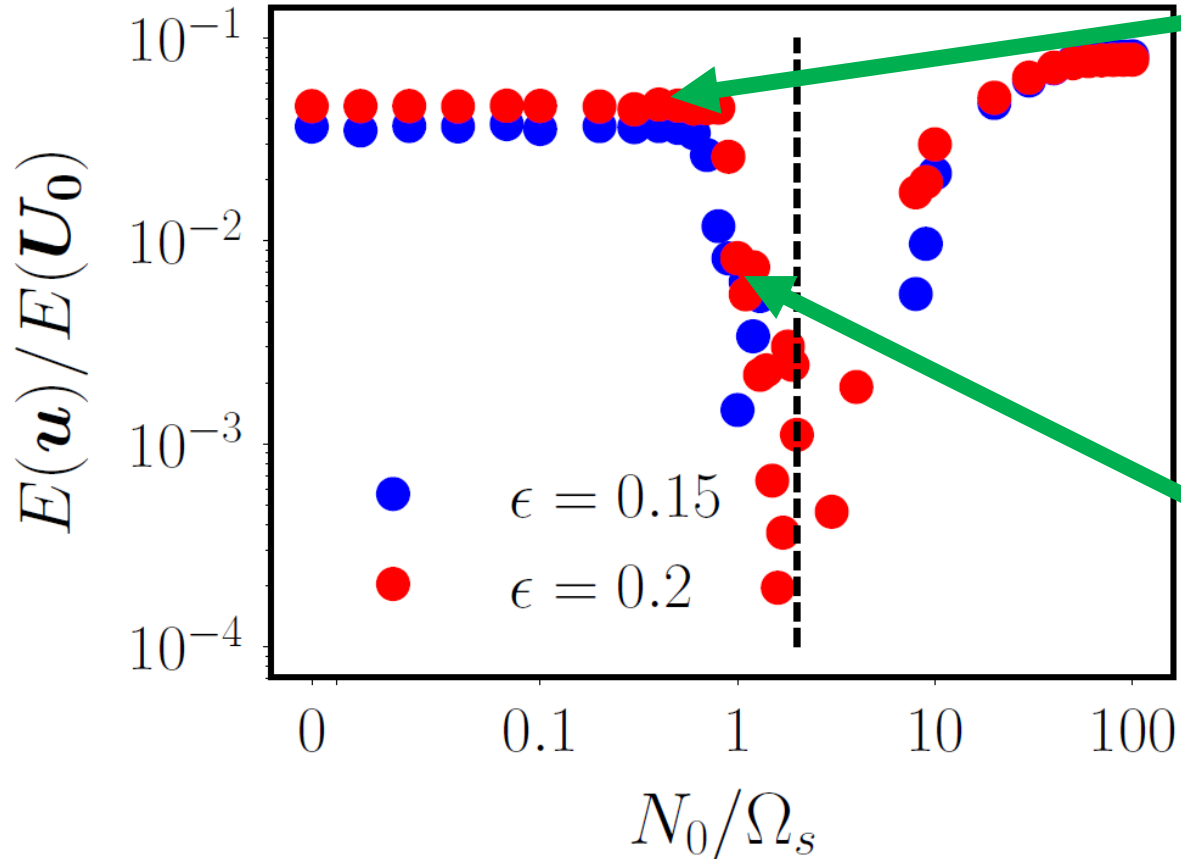
Total temperature

$\epsilon = 0.2$
 $N_0/\Omega_s = 0.5$

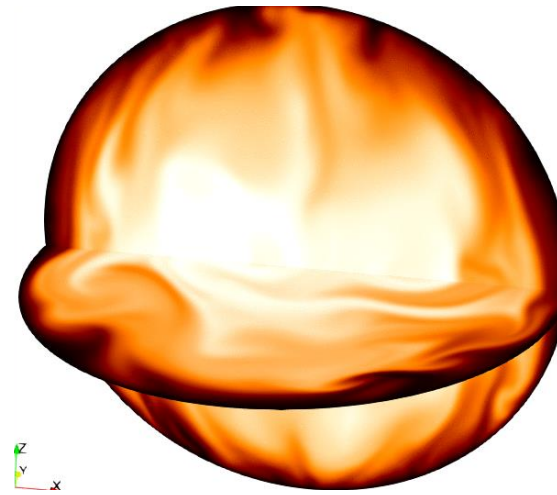
Hydrodynamic instability

(Vidal et al. MNRAS, 2018)

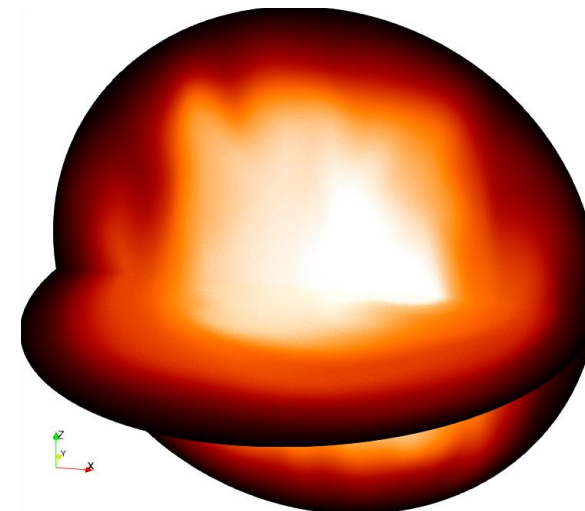
- ▶ $Ek = 10^{-4}$, $Pr = 1$, $\Omega_0 = 0$,
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Total temperature



$\epsilon = 0.2$
 $N_0/\Omega_s = 0.5$



$\epsilon = 0.2$
 $N_0/\Omega_s = 1$

MHD simulations

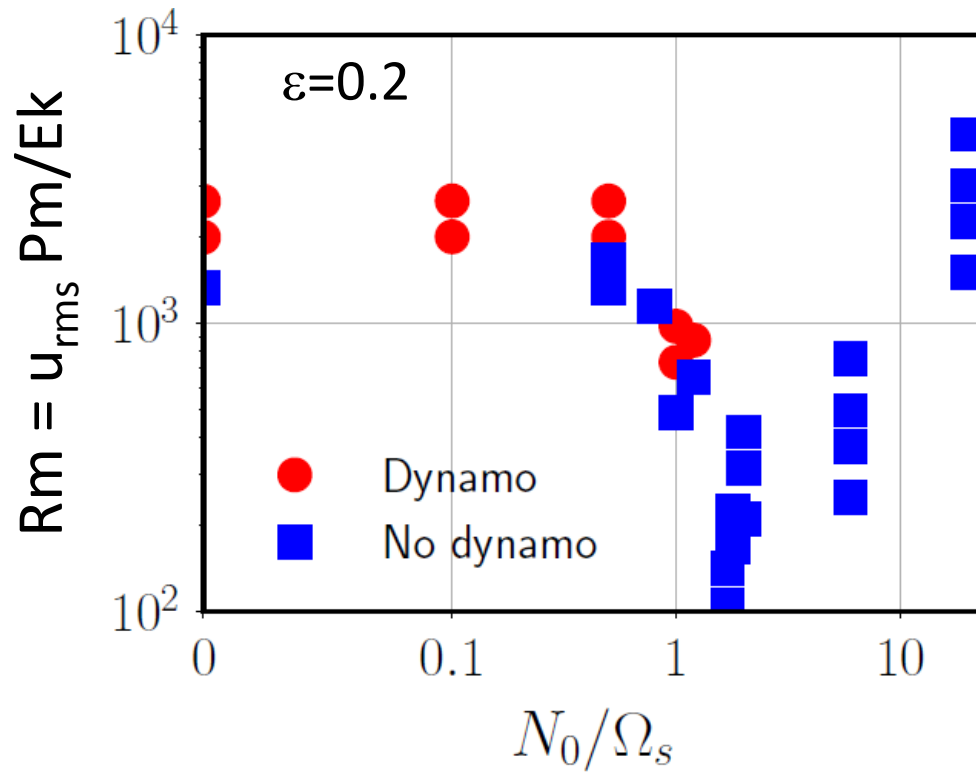
(Vidal et al., MNRAS, 2018)

- ▶ Basic flow U_0 **not dynamo capable** (for magnetic Prandtl numbers $Pm \leq 5$),
- ▶ Integration over one **magnetic diffusive time**.

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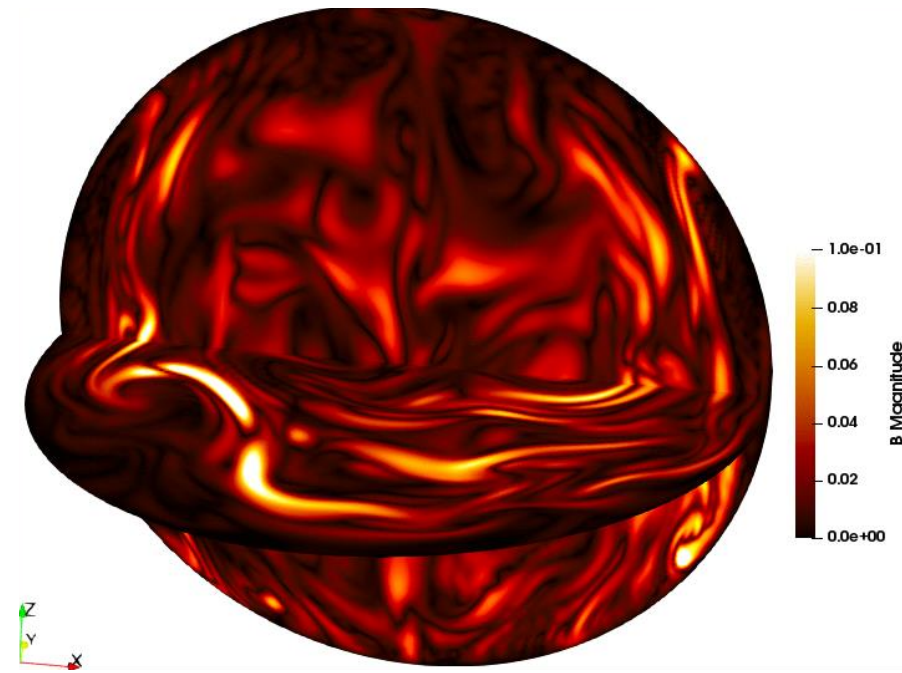
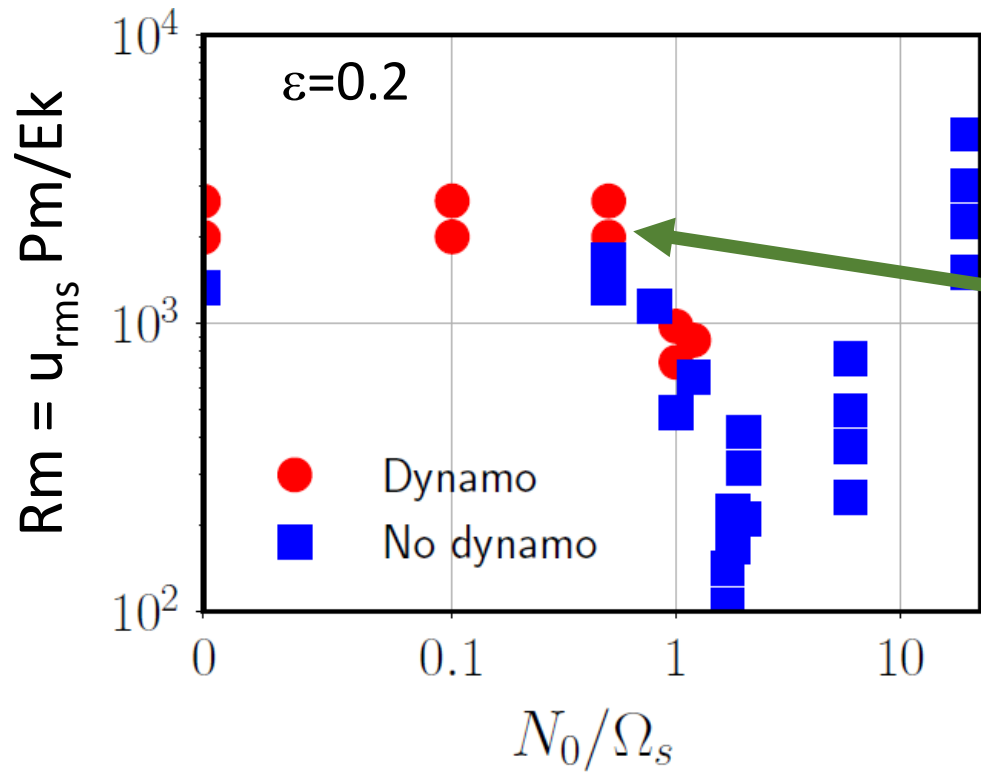


Dynamo for $Rm > 2000$ & $Pm \sim 1$

MHD simulations

(Vidal et al., MNRAS, 2018)

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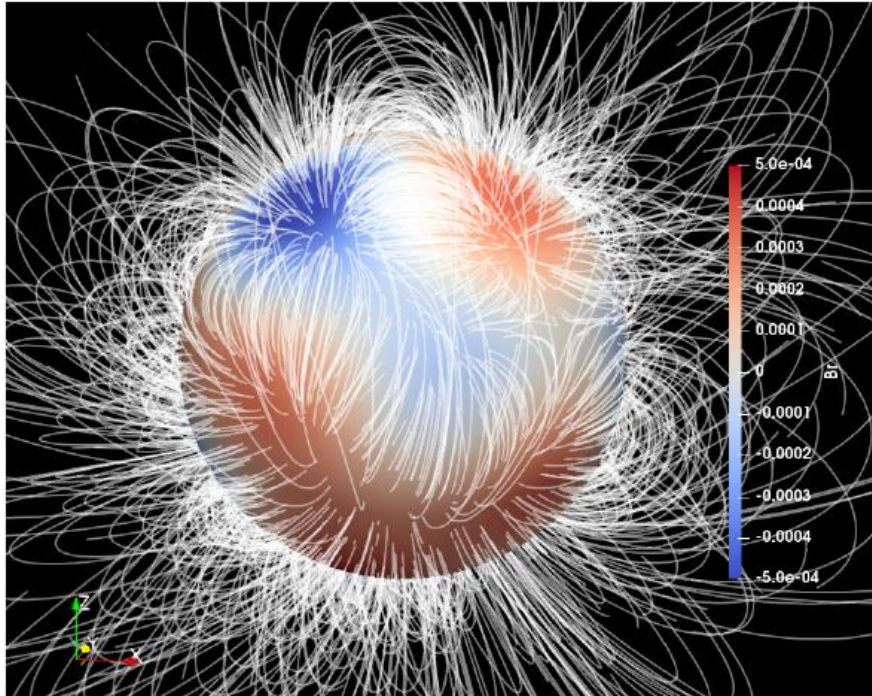


Dynamo for $Rm > 2000$ & $Pm > 1$

Saturated dynamo

Extrapolation

(Vidal et al., MNRAS, 2018)



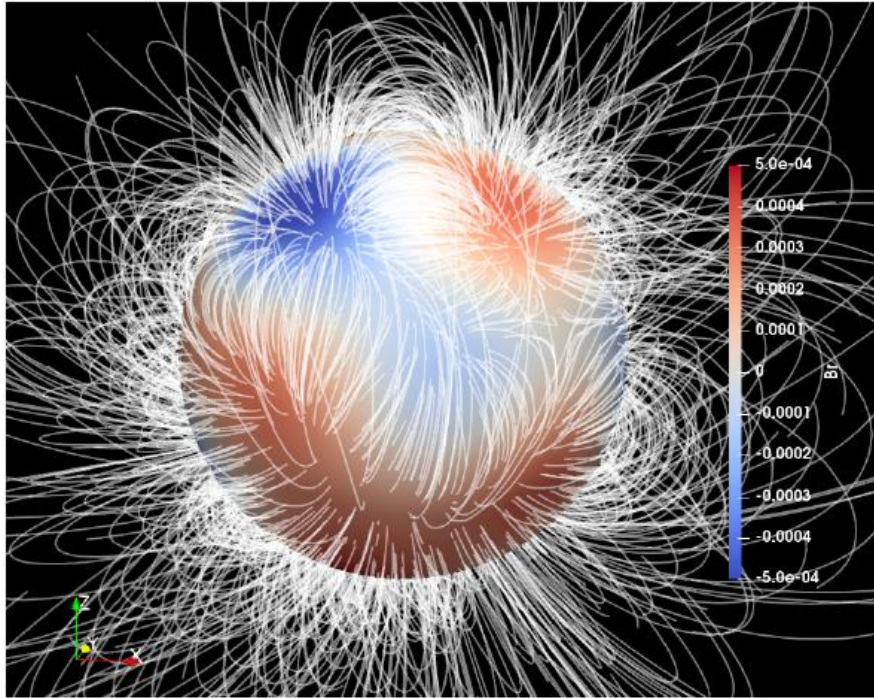
Large-scale surface field!

- ▶ Similar to convective scalings,
- ▶ Typical surface field **strength**

$$B_0 = \delta \frac{3}{2} \sqrt{\frac{3\mu_0}{4\pi}} \frac{R_*^{5/2}}{M_*^{1/2}} \Omega_s \frac{m}{D^3} |1 - \Omega_0|, \quad 10^{-3} \leq \delta \leq 10^{-2}$$

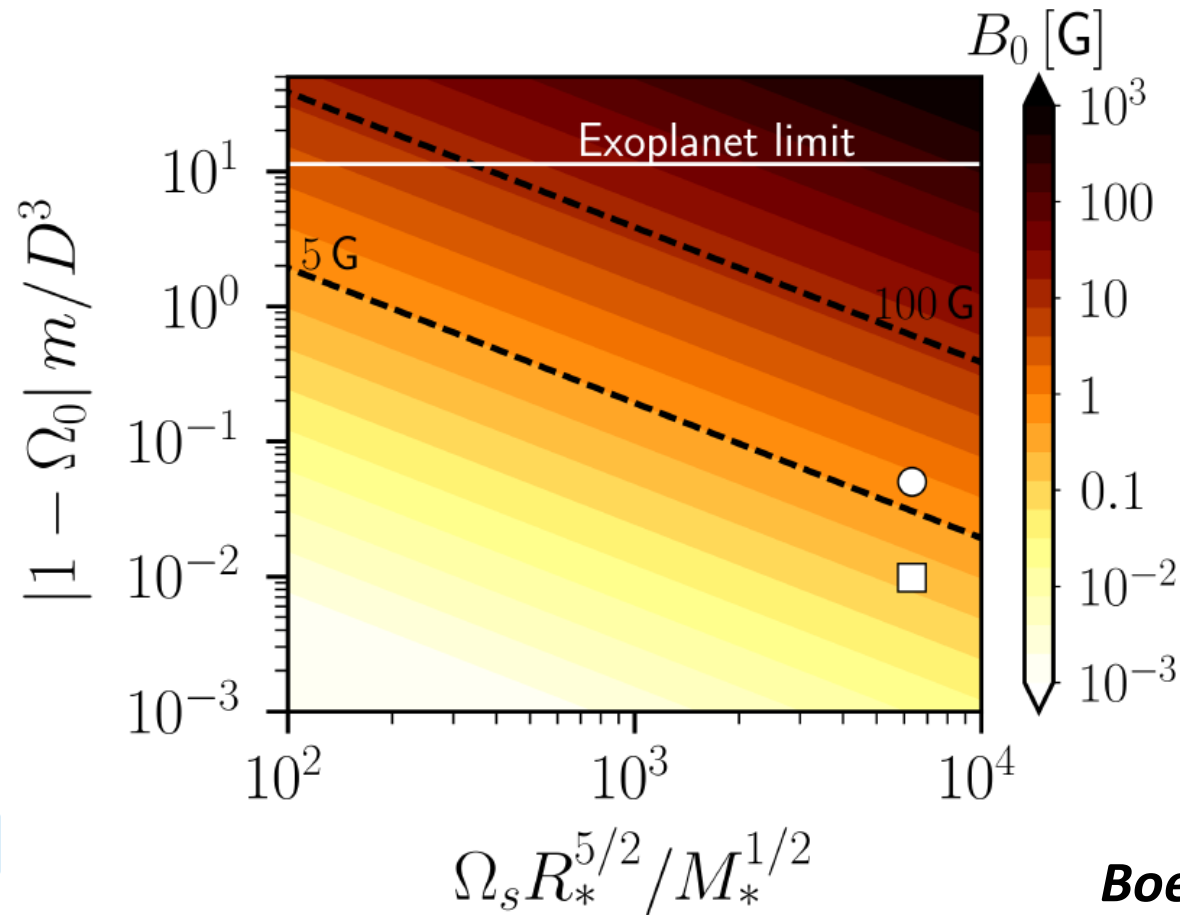
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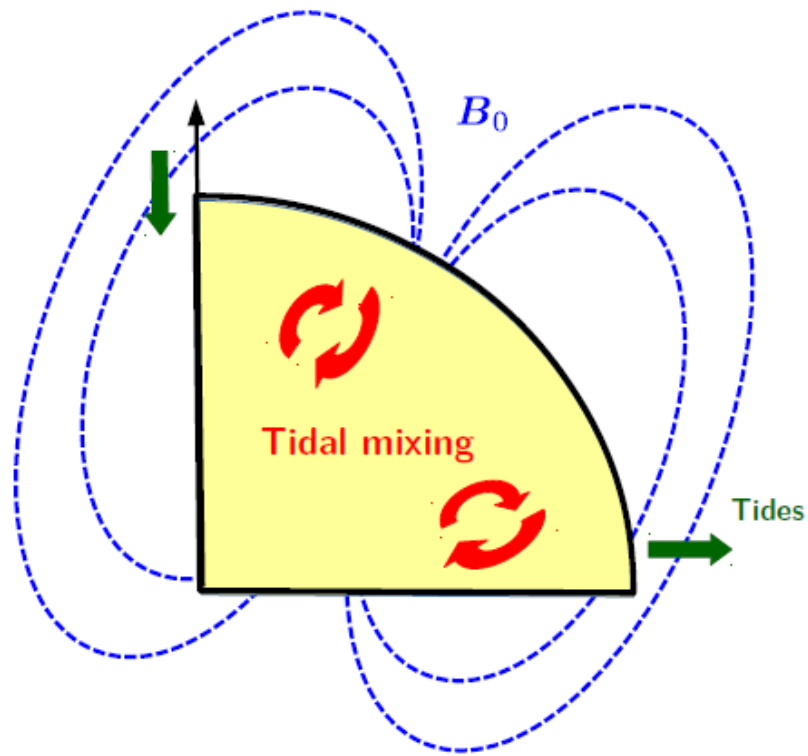
Vega
Measure:
 0.6 ± 0.3 G
Theory:
1 - 1.5 G

Boehm et al., 2015

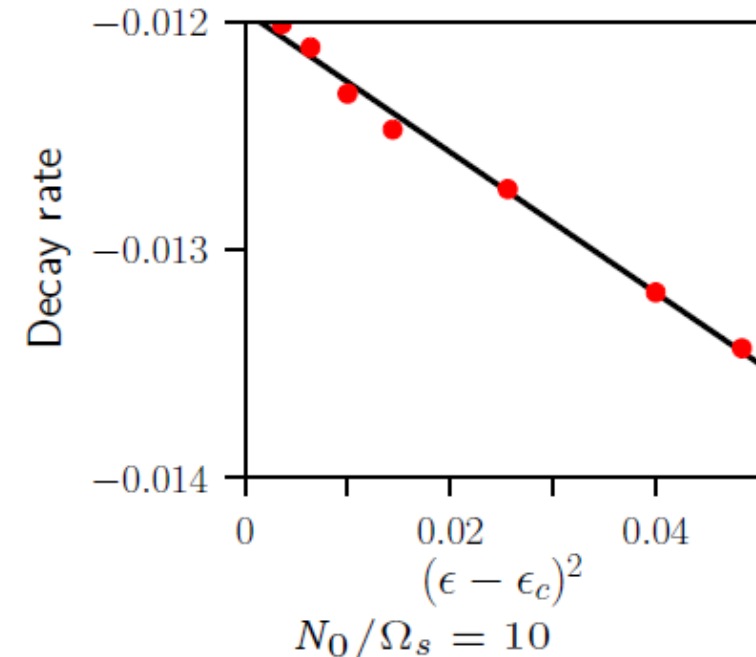
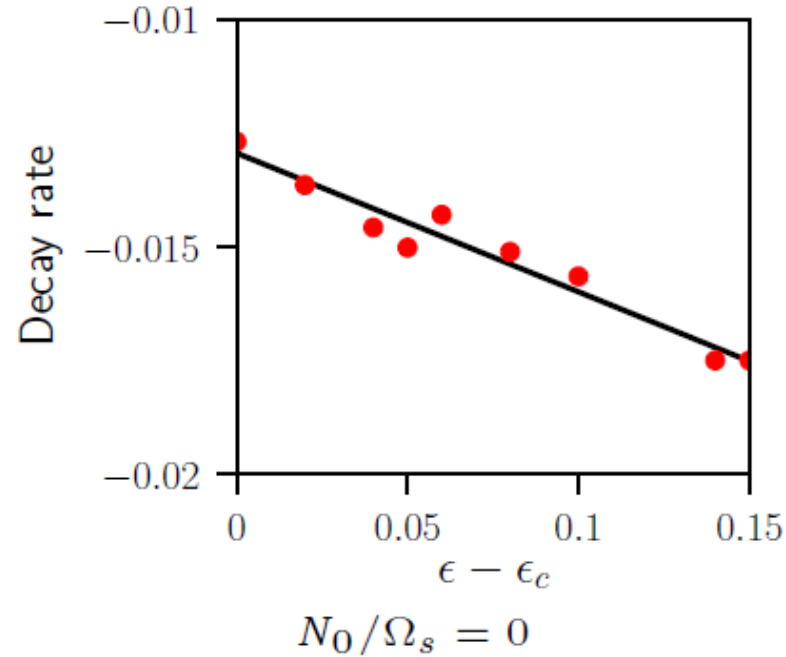
$$B_0 = \delta \frac{3}{2} \sqrt{\frac{3\mu_0}{4\pi}} \frac{R_*^{5/2}}{M_*^{1/2}} \Omega_s \frac{m}{D^3} |1 - \Omega_0|$$

Perspective: tidal mixing of fossil fields

Ongoing collaboration with Dr. Evelyne Alecian (IPAG) & Dr. Asif ud-Doula (Dunmore)



Enhancing of Ohmic decay!



Conclusion

1) Tides can lead to mixing & dynamos in stratified fluids

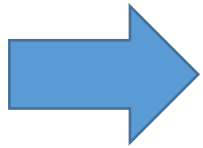
- For moderate stratification ($N/\Omega < 10$, see Vidal+18)
- For large enough tides

2) Tides can accelerate the fossile field (Ohmic) decay via the mixing

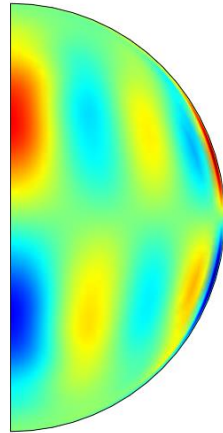
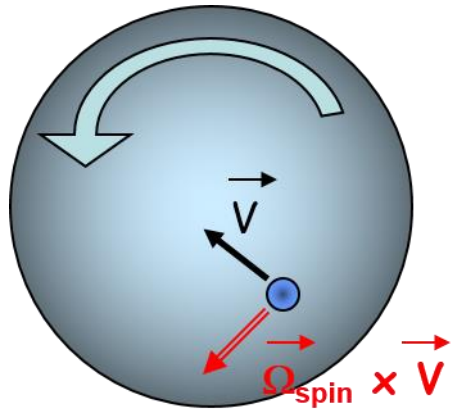
Thank you!

Inertial instabilities

Restoring force
(Coriolis, buoyancy,
Lorentz force, etc.)

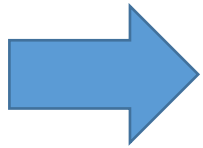


Waves, modes
(inertial, internal,
Alfven, etc.)



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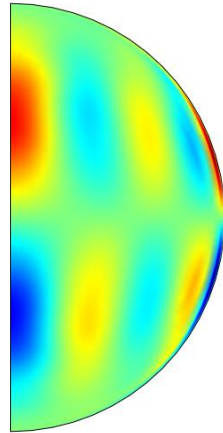
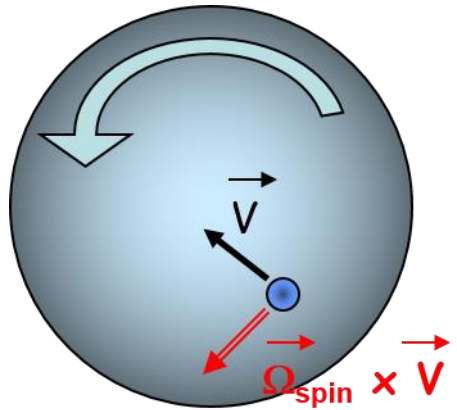
Waves, modes
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Periodic modulation

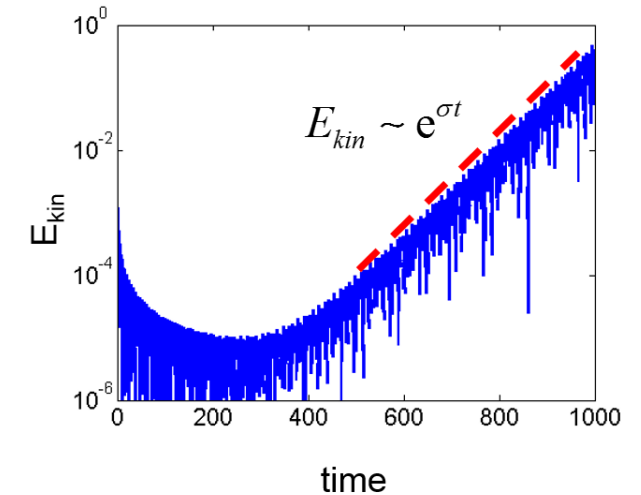


e.g. tides, precession, etc.

Parametric resonance
(tidal or precession instability,
Faraday instability, etc.)



$$\sim \cos[m \theta(t)]$$



Precession topographic instability ~ **Tidal instability** : similar mechanism
=> **inertial instability** (parametric resonance)